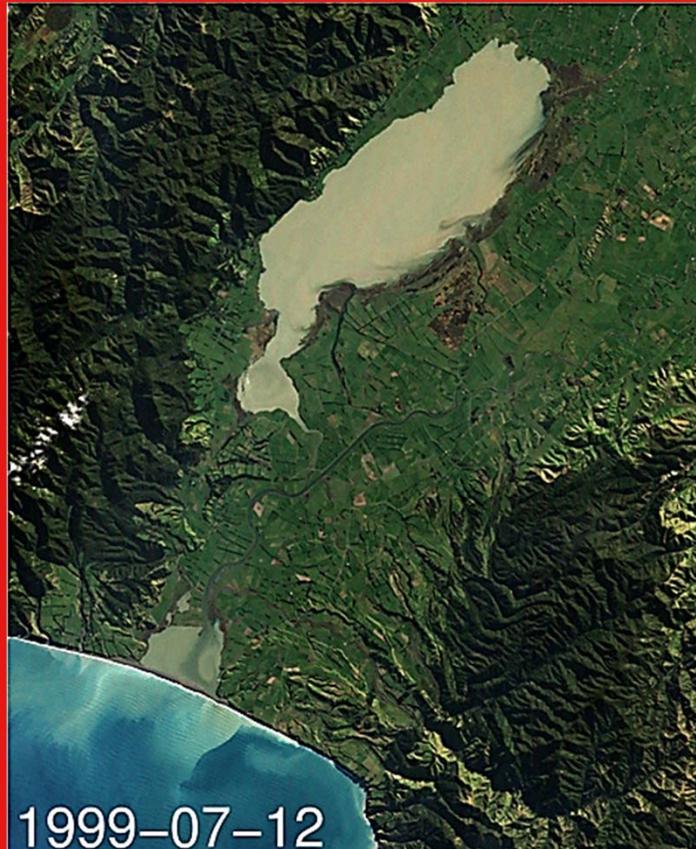




THE UNIVERSITY OF  
**WAIKATO**  
*Te Whare Wānanga o Waikato*

## *Modelling and remote sensing of lakes Onoke and Wairarapa*

*Mathew Allan*  
*David Hamilton*



2016

## Approaches to assess water quality

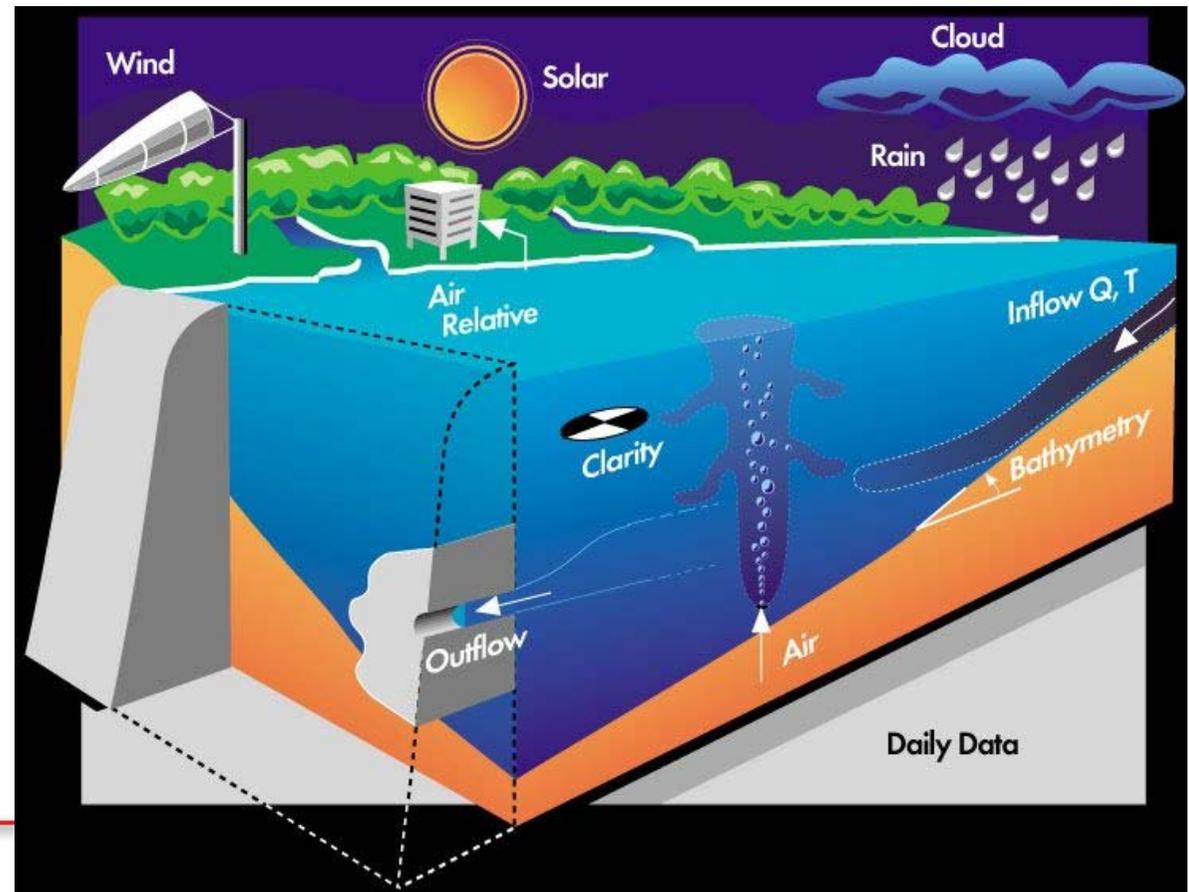
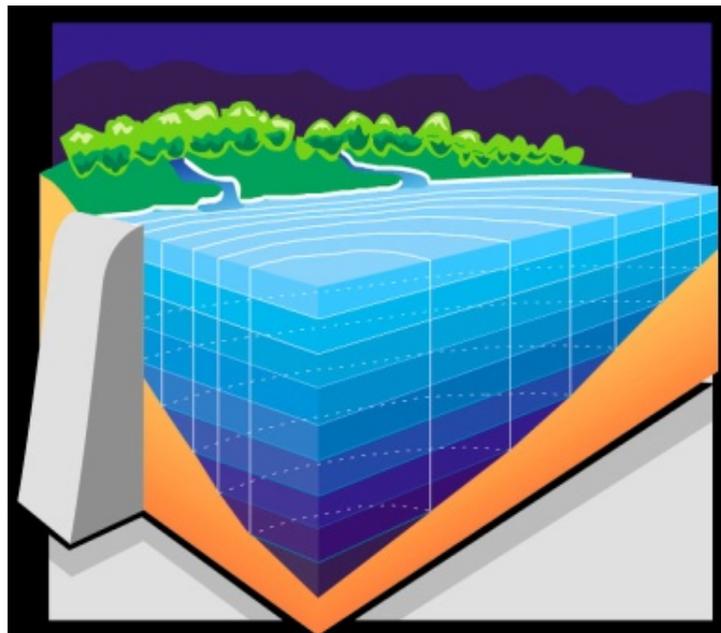
- In situ sampling, numerical modelling and remote sensing
- In situ sampling lacks the ability to effectively monitor the spatial variability
- Three-dimensional (3-D) and one-dimensional (1-D) hydrodynamic modelling of lake water quality and temperature:
  - interpolate temporal gaps in data derived from satellite and traditional monitoring
  - extend the analysis to the vertical domain
  - provide insights into the spatial variability of biogeochemical processes

# Method: 3-D Numerical Hydrodynamic Model

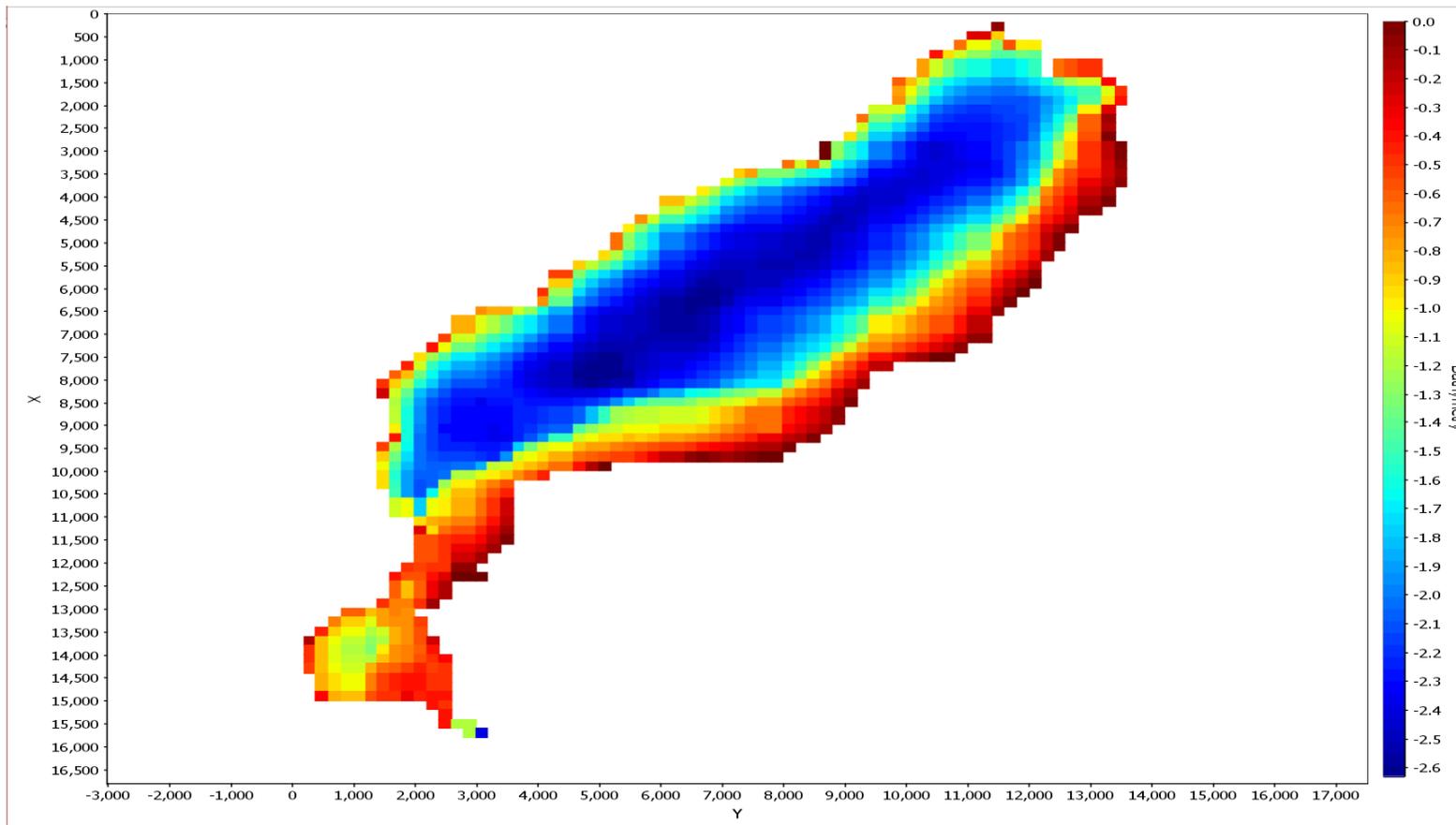
The Estuary and Lake Computer Model (ELCOM)

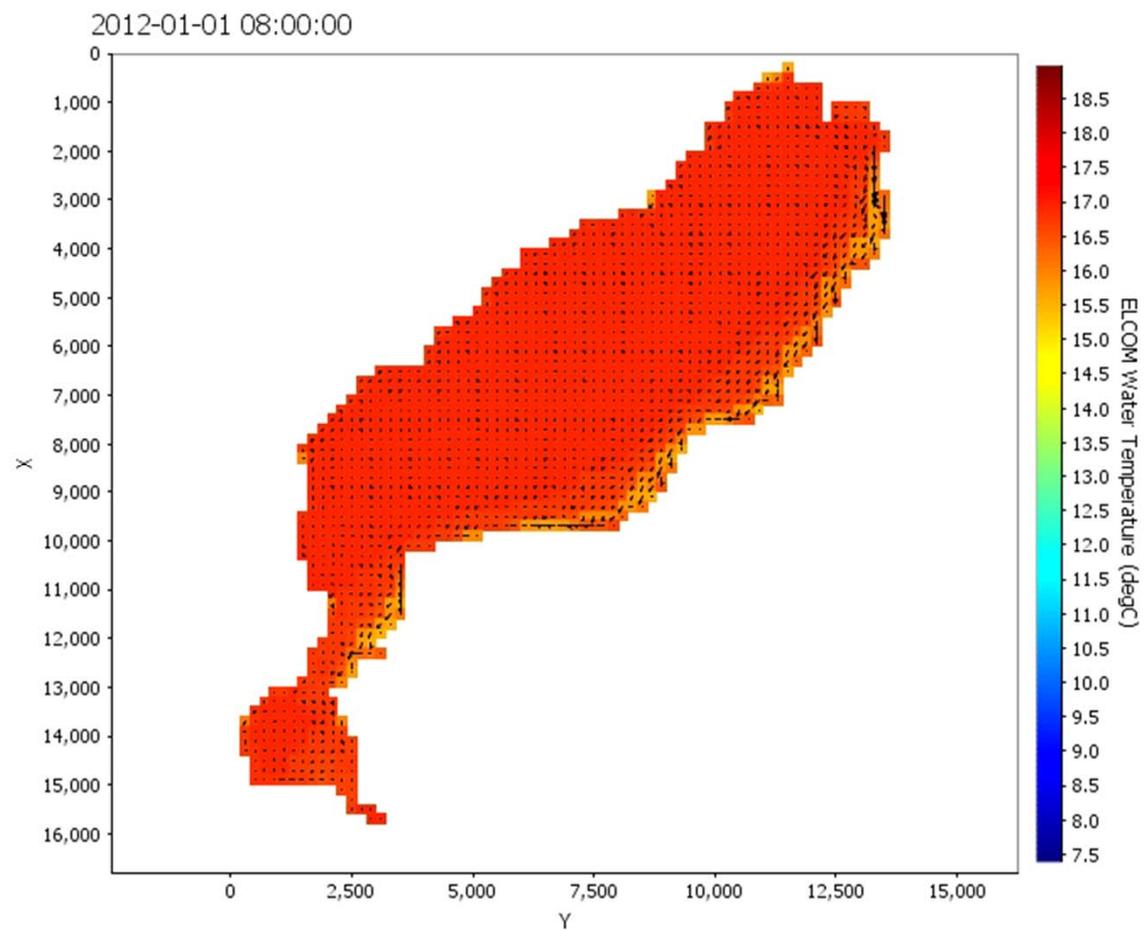
CWR Australia

Hydrodynamic and thermodynamic models in order to predict velocity, salinity and temperature in waterbodies

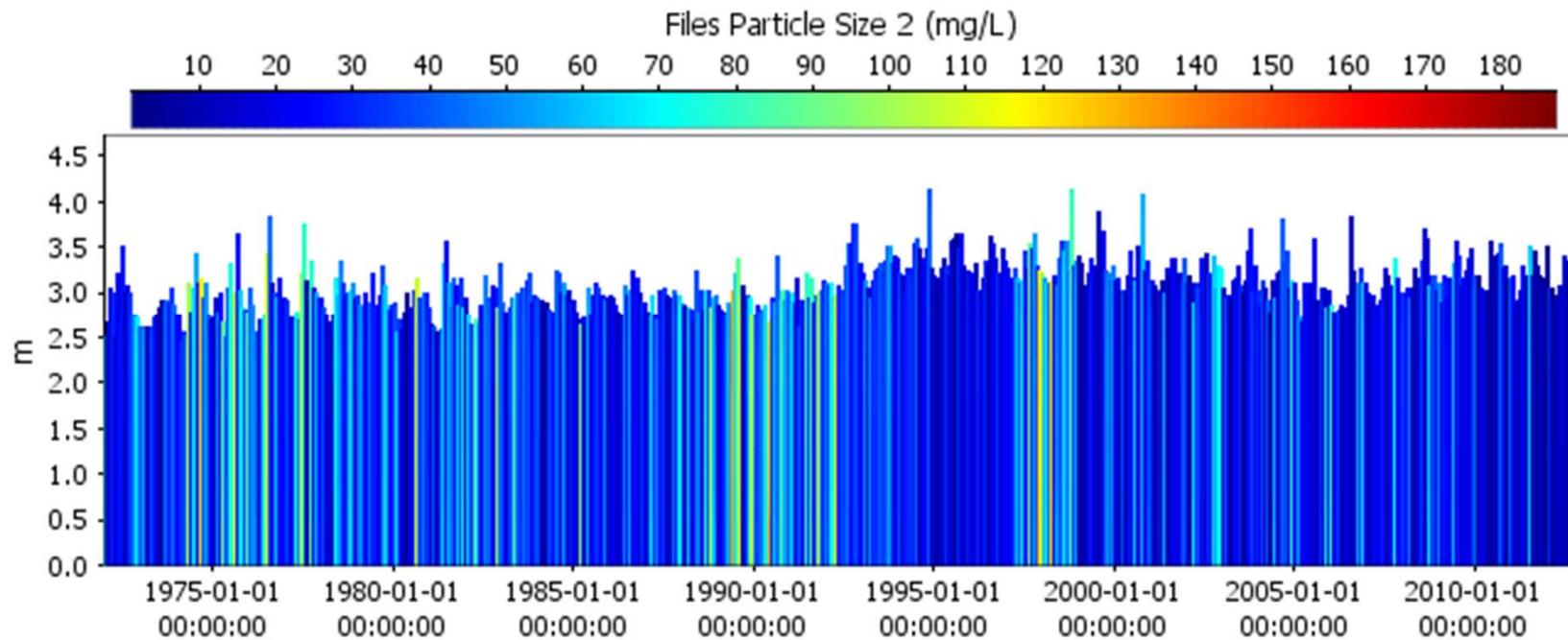


# 3-D model





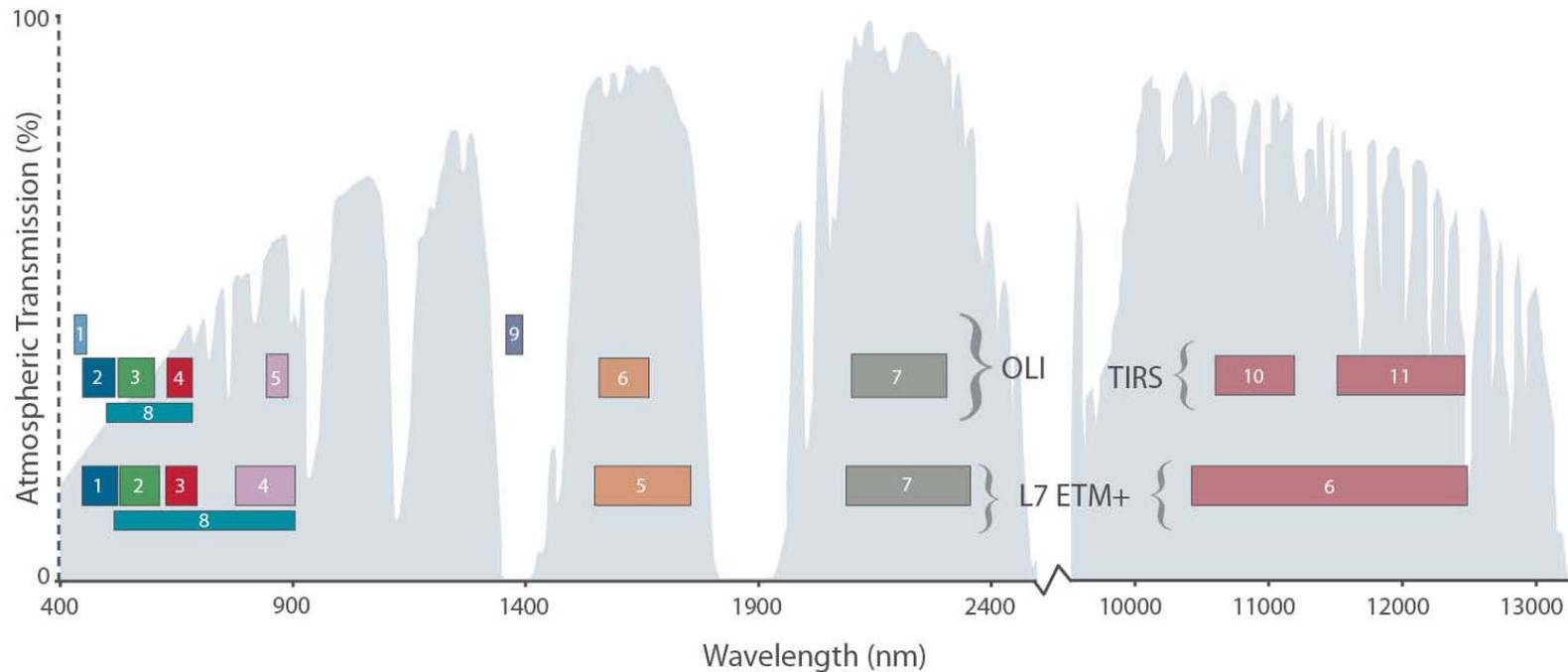
# 1-D model



# *Modelling- future work*

- Calibration/validation of 1D and 3-D models using in situ water quality data measured by the Greater Wellington Regional Council and remote sensing data
- Calibrated models will then be used to determine the likely water quality and ecological effects of specific management scenarios

# Remote sensing data



United States Geological Survey (USGS) on demand atmospherically corrected Landsat imagery ordered from <http://espa.cr.usgs.gov/index/>, including 91 images

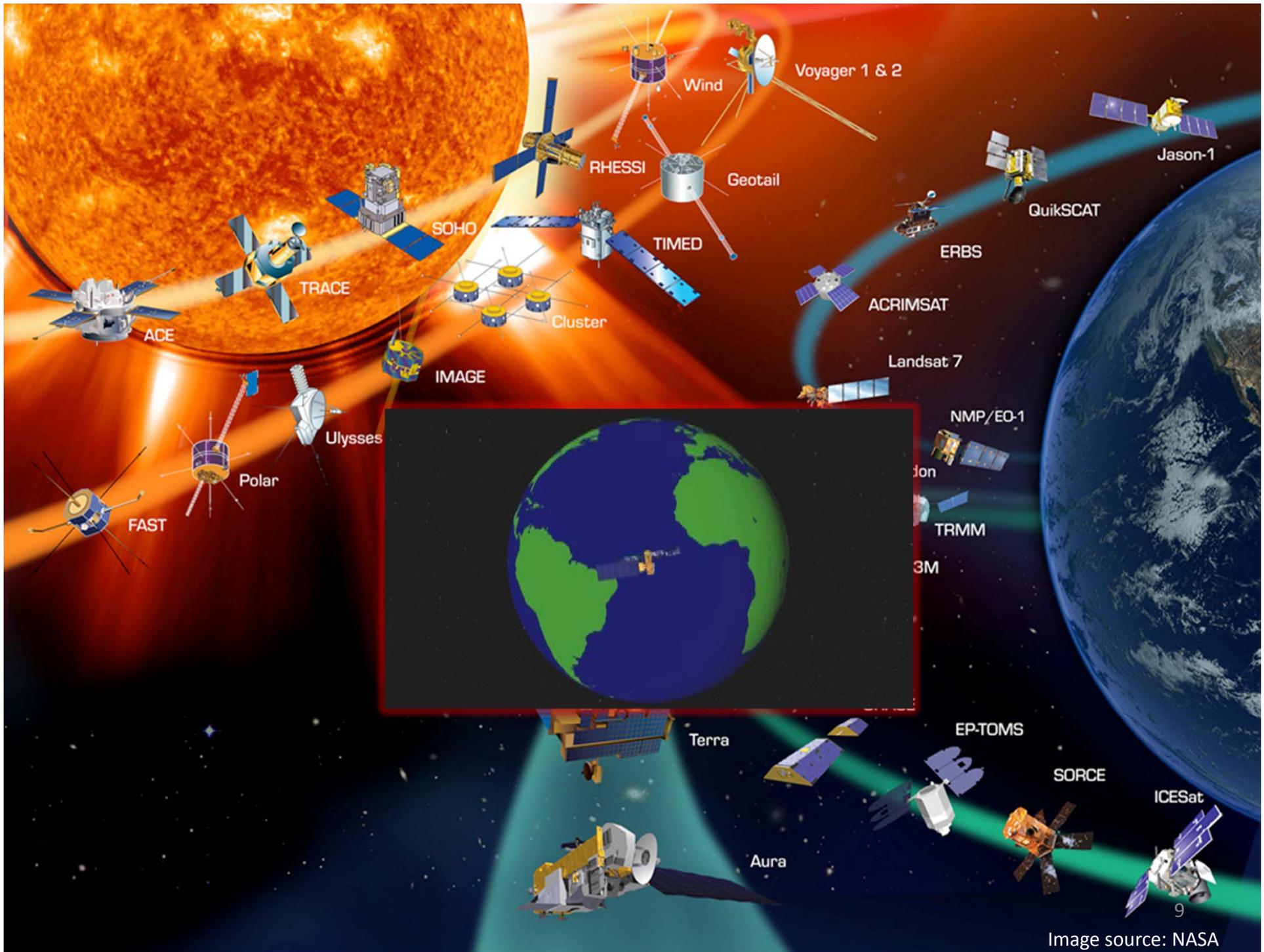


Image source: NASA

# Remote sensing semi-analytical model

Incident Light from the Sun

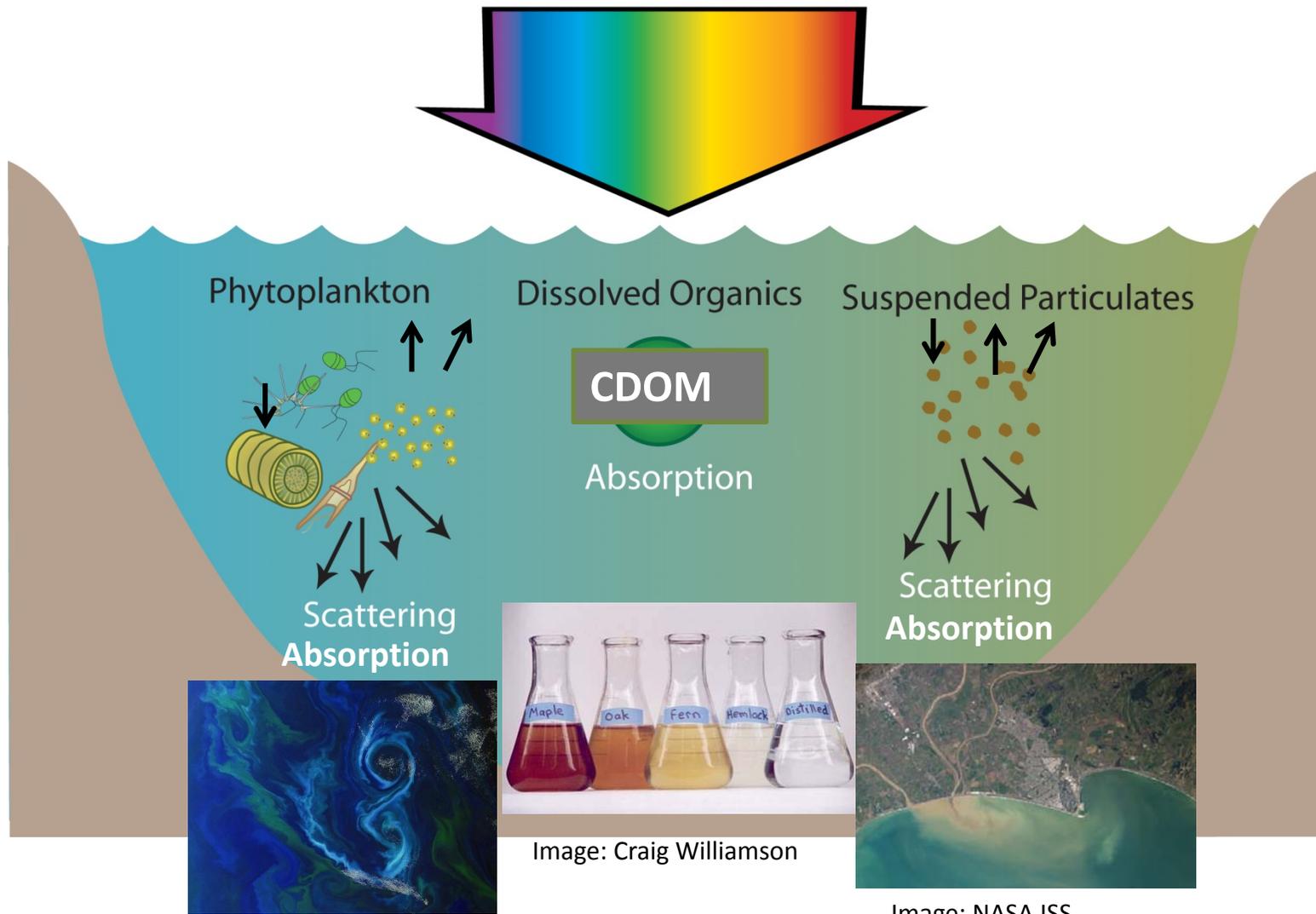


Image: NASA

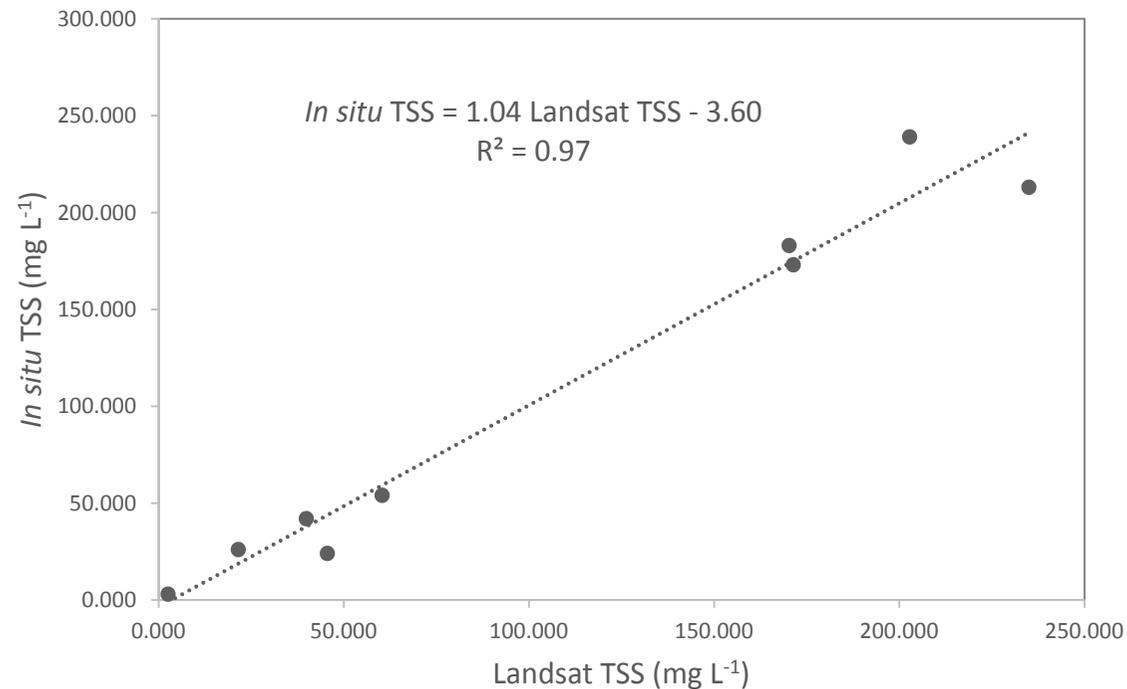


Image: Craig Williamson



Image: NASA ISS

# Remote sensing results



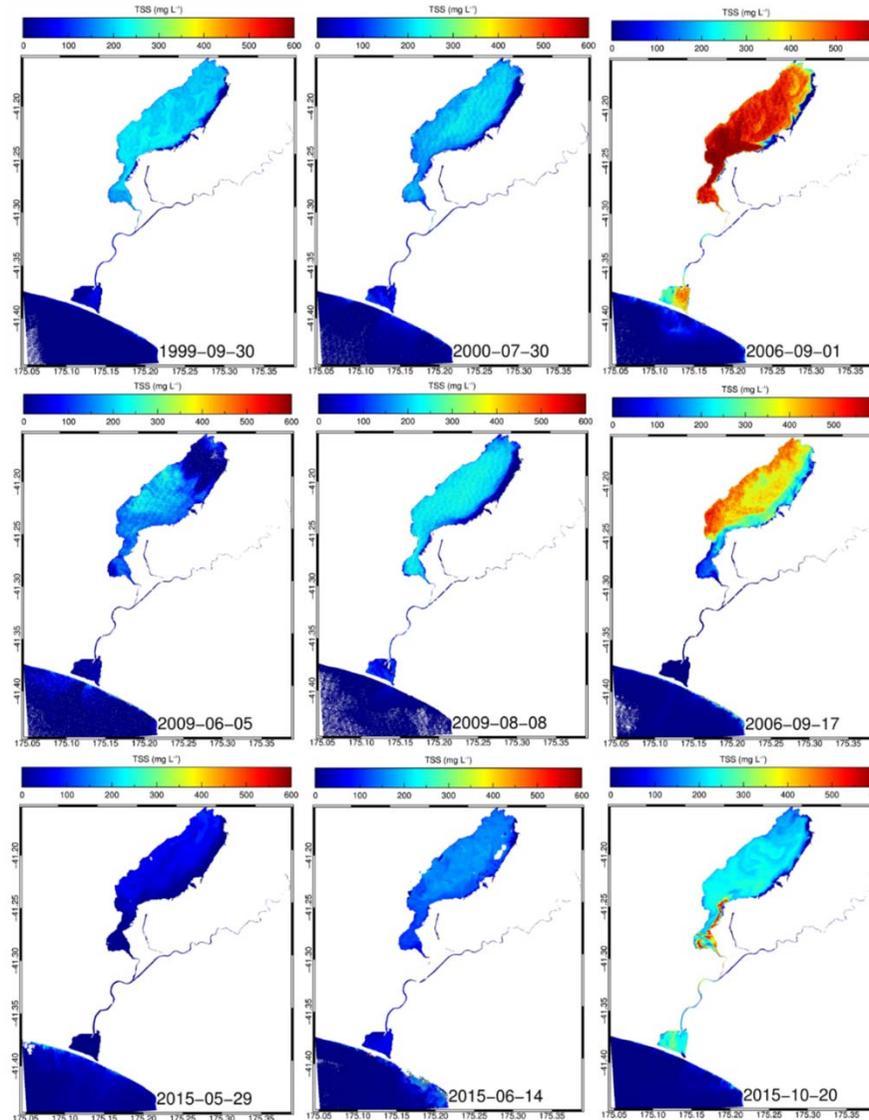
**Comparison of *in situ* TSS to Landsat estimated TSS (mg L<sup>-1</sup>) using a semi-analytical relationship**

$r^2$  of 0.97,

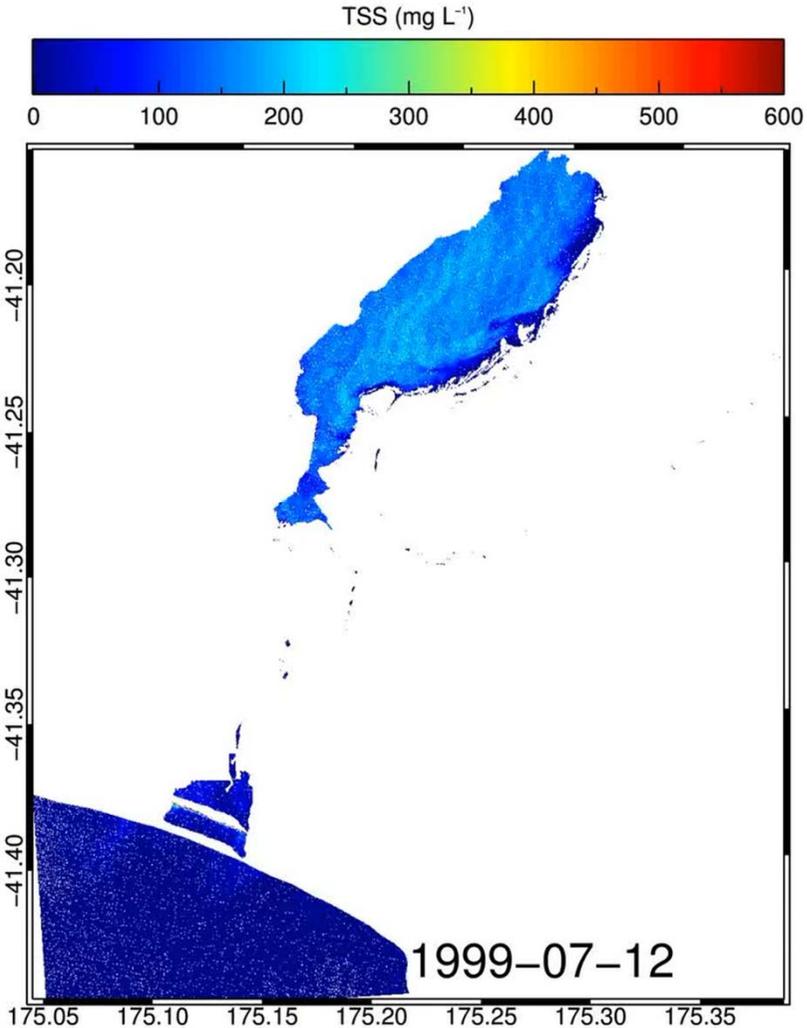
RMSE 16.62 mg L<sup>-1</sup>

32% RMSE over a range of *in situ* TSS from 3 - 239 mg L<sup>-1</sup>.

# Landsat estimated total suspended solids

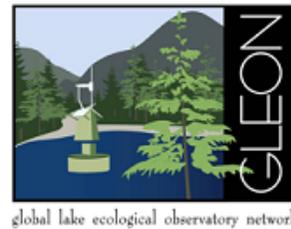


# Landsat estimated total suspended solids

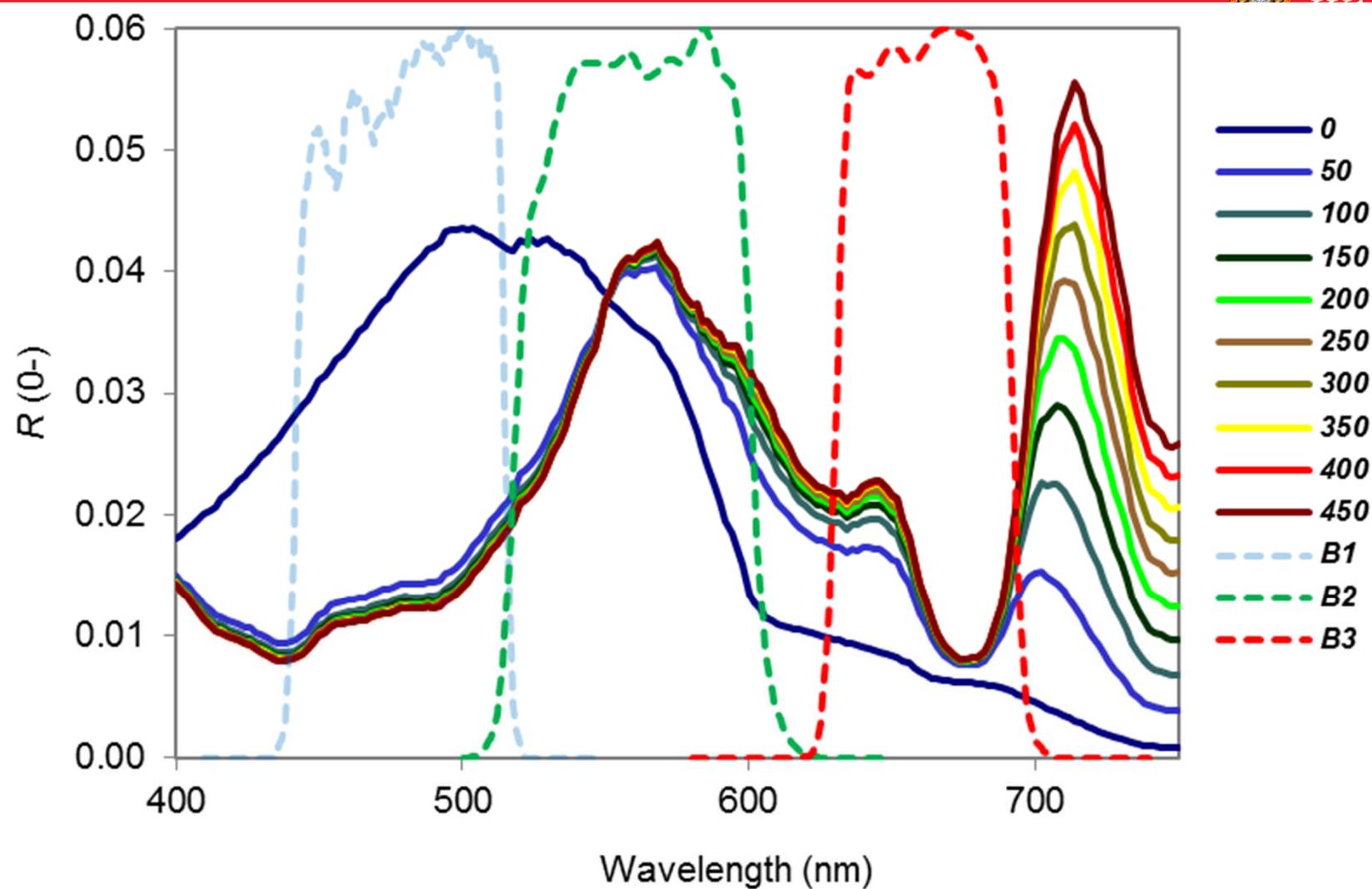


# Acknowledgements

- Greater Wellington Regional Council staff
- Ruamāhanga Whaitua Committee
- Chris McBride (UOW)

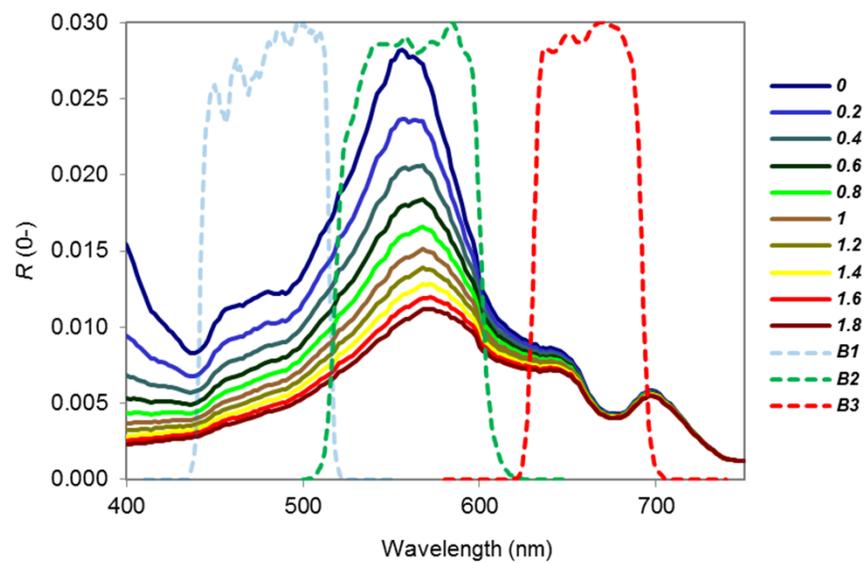


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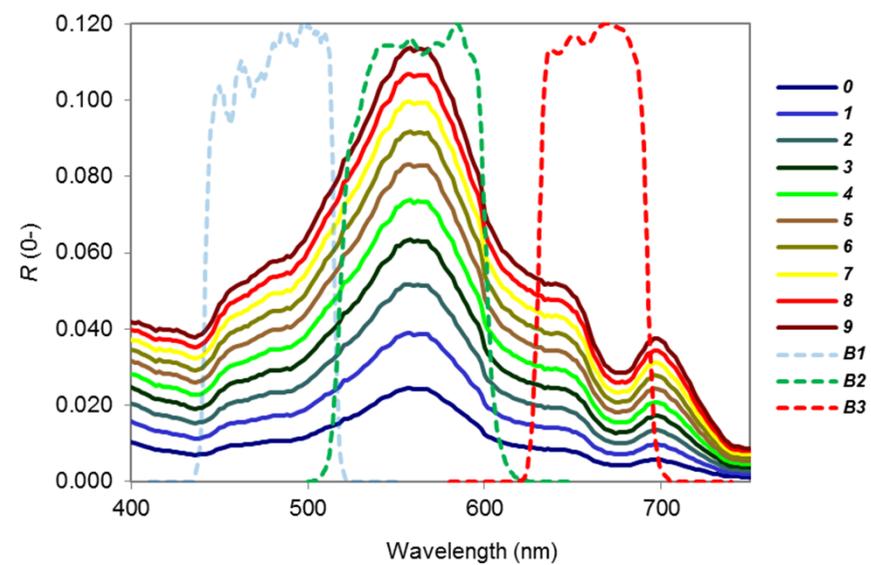


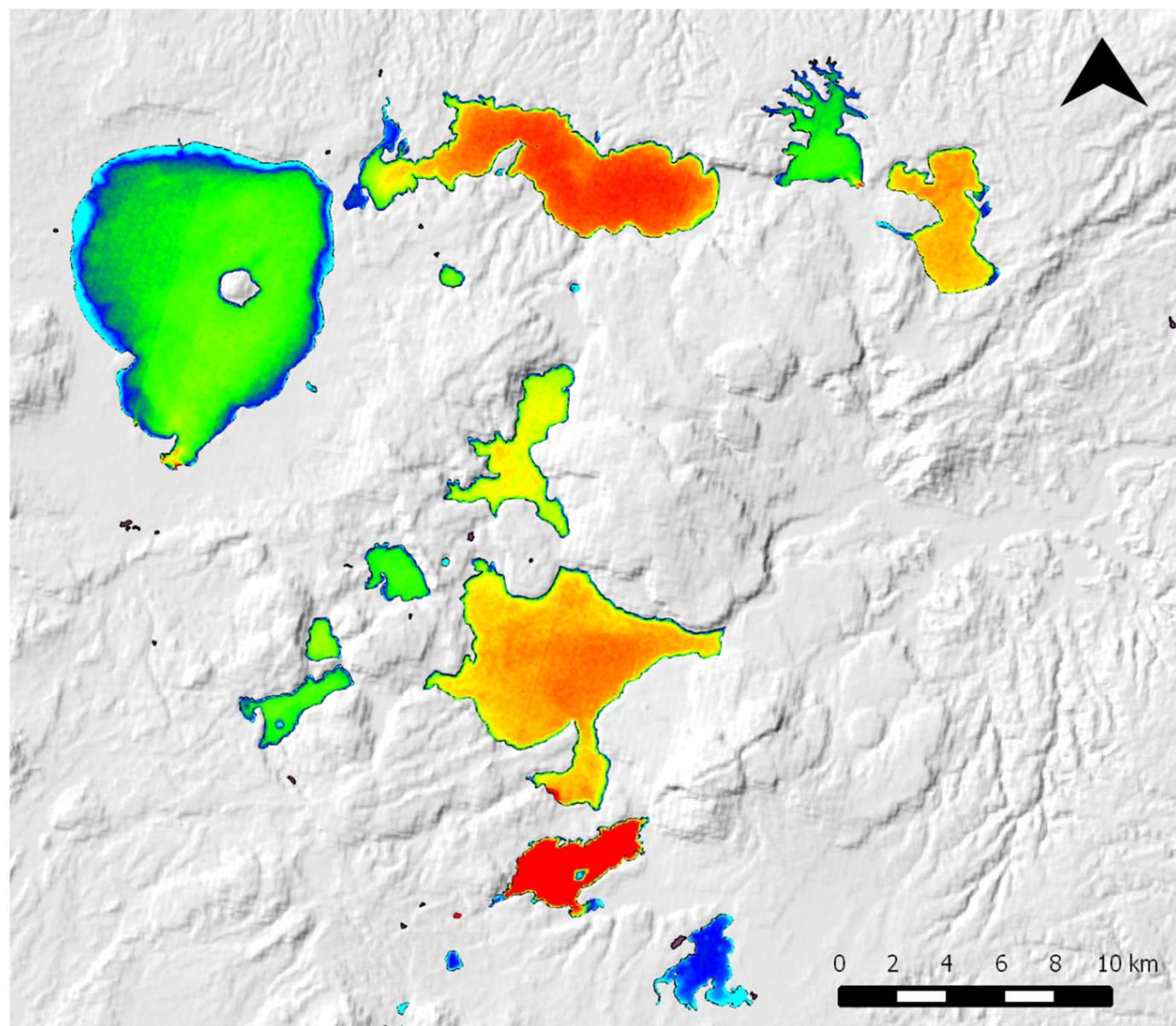
Modelled subsurface reflectance ( $R(0-)$ ) for varying chlorophyll *a* concentrations ranging from 0 to 450  $\mu\text{g L}^{-1}$  with fixed CDOM absorption of  $0.16 \text{ m}^{-1}$  and tripton concentration of  $0.5 \text{ mg L}^{-1}$ . The relative spectral response of Landsat bands B1, B2 and B3 is overlaid.

## CDOM



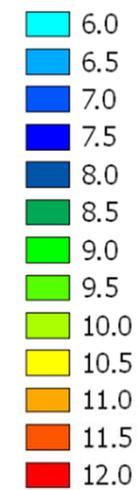
## TSS – non-living





Temperature (°C)

24-06-2013



# Equations RS

$$r_{rs}(\lambda) = g_0 u(\lambda) + g_1 [u(\lambda)]^2$$
$$u(\lambda) = \frac{b_b(\lambda)}{a(\lambda) + b_b(\lambda)}$$

- 
- 
- and  $g_0$  and  $g_1$  are empirical constants that depend on the anisotropy of the downwelling light field and scattering processes within the water. The constant  $g_0$  is equivalent to  $f/Q$  where  $f$  represents geometrical light factors and  $Q$  represents the light distribution factor, which is defined as upwelling subsurface irradiance/upwelling subsurface radiance (Maritorena et al. 2002). It has been suggested that  $g_0$  and  $g_1$  should be considered as variables in optically complex inland waters (Aurin and Dierssen 2012; Li et al. 2013) therefore we used fitted values for  $g_0$  and  $g_1$  of 0.103 and 0.009 respectively, which were derived previously for Lake Ellesmere (Allan, 2014).
- The absorption and backscattering coefficients are comprised of individual optically active constituents:
  - $b_b(\lambda) = bbw(\lambda) + BpTSS \cdot b^*TSS(\lambda) \cdot CTSS$
  - $a(\lambda) = aw(\lambda) + C\varphi \cdot a^*\varphi(\lambda) + aCDOMD(\lambda)$
  - $aCDOMD(\lambda) = aCDOMD(\lambda_{440}) \exp[-S(\lambda - \lambda_{440})]$
  - where:
    - $bbw(\lambda)$  = backscattering coefficient of water
    - $BpTSS$  = backscattering ratio from TSS
    - $b^*TSS(\lambda)$  = specific scattering coefficient of TSS
    - $CTSS$  = concentration of TSS
    - $b^*\varphi(\lambda)$  = specific scattering coefficient of phytoplankton
    - $aw(\lambda)$  = absorption coefficient of pure water
    - $C\varphi$  = concentration of chl  $a$
    - $a^*\varphi(\lambda)$  = specific absorption coefficient of phytoplankton
    - $aCDOMD(\lambda)$  = absorption coefficient for coloured dissolved organic matter (CDOM)
    - $S$  = spectral slope coefficient
  - Values of  $aw(\lambda)$  and  $bbw(\lambda)$  were prescribed from the literature (Morel 1974; Pope and Fry 1997). The backscattering ratio of TSS,  $BpTSS$ , was set to 0.019 (Petzold 1972). The specific scattering coefficient of TSS at the Landsat b3 wavelength was estimated using a power function (Morel and Prieur 1977) where the value  $b^*TSS(555)$  was set to 0.6  $m^2 \cdot g^{-1}$ . The hyperbolic exponent  $n$  was set to 0.63, equating to a value measured in Lake Taupo, New Zealand (Belzile et al. 2004). The  $a^*\varphi(662)$  was 0.0136  $m^2 \cdot mg^{-1}$ , equal to the average value measured in eight Dutch lakes (Dekker et al., 2002). The bio-optical simulations were run by varying TSS concentration from 0.1 to 417.6  $mg \cdot L^{-1}$  in increments of 0.5  $mg \cdot L^{-1}$  while  $aCDOMD(440)$  was fixed at 0.042  $m^{-1}$ , which is average in situ of measurement in Lakes Wairarapa and Onoke, with chl  $a$  ( $\mu g \cdot L^{-1}$ ) taken to increase with TSS (chl  $a = TSS/6$ ) ranging from 0.017 to 67.6  $\mu g \cdot L^{-1}$  (encompassing a similar range to that measured in situ in Lakes Onoke and Wairarapa).